
Article:
Hughes, Claire orcid.org/0000-0002-9512-8052, Barton, Eleanor, Hepach, Helmke et al. (7 more authors) (2021) Oxidation of iodide to iodate by cultures of marine ammonia-oxidising bacteria. MARINE CHEMISTRY. 104000. ISSN 0304-4203

https://doi.org/10.1002/essoar.10503313.2

Oxidation of iodide to iodate by cultures of marine ammonia-oxidising bacteria

Claire Hughes[1*], Eleanor Barton[1*], Helmke Hepach[1], Rosie Chance[2], Matt Pickering[1], Karen Hogg[3], Andreas Pommerening-Röser[4], Martin R. Wadley[5], David P. Stevens[5] and Tim D. Jickells[6]

[1]Department of Environment and Geography, University of York, Wentworth Way, Heslington, York, YO10 5NG, UK

[2]Wolfson Atmospheric Chemistry Laboratory, Department of Chemistry, University of York, Heslington, York, YO10 5DD, UK

[3]Bioscience Technology Facility, Department of Biology, University of York, Wentworth Way, York, YO10 5DD, UK

[4]University of Hamburg, Mikrobiologie & Biotechnologie, Ohnhorststr. 18, D-22609 Hamburg, Germany

[5]School of Mathematics, University of East Anglia, Norwich Research Park, Norwich, NR4 7TJ, UK

[6]School of Environmental Sciences, University of East Anglia, Norwich Research Park, Norwich, NR4 7TJ, UK

*authors contributed equally to the manuscript

Corresponding author: Rosie Chance (rosie.chance@york.ac.uk)

Highlights:

- Oxidation of iodide to iodate by marine nitrifying bacteria demonstrated for first time
- Laboratory cultures of ammonium oxidising bacteria produced iodate from iodide substrate
- Nitrification used to parameterise iodide sink in global marine iodine cycling model
- Changes in nitrification may increase sea surface iodide, impacting atmospheric chemistry

Abstract

Reaction with iodide (I^-) at the sea surface is an important sink for atmospheric ozone, and causes sea-air emission of reactive iodine which in turn drives further ozone destruction. To incorporate this process into chemical transport models, improved understanding of the factors controlling marine iodine speciation, and especially sea-surface iodide concentrations, is needed. The oxidation of I^- to iodate (IO_3^-) is the main sink for oceanic I^-, but the mechanism for this remains unknown. We demonstrate for the first time that marine nitrifying bacteria mediate I^- oxidation to IO_3^-. A significant increase in IO_3^- concentrations compared to media-only controls was observed in cultures of the ammonia-oxidising bacteria *Nitrosomonas* sp. (Nm51) and *Nitrosoccocus oceani* (Nc10) supplied with 9-10 mM I^-, indicating I^- oxidation to IO_3^-. Cell-normalised production rates were 15.69 (±4.71) fmol IO_3^- $cell^{-1}$ d^{-1} for *Nitrosomonas* sp., and 11.96 (±6.96) fmol IO_3^- $cell^{-1}$ d^{-1} for *Nitrosococcus oceani*, and molar ratios of iodate-to-nitrite production were 9.2±4.1 and 1.88±0.91 respectively. Preliminary experiments on nitrite-oxidising bacteria showed no evidence of I^- to IO_3^- oxidation. If the link between ammonia and I^- oxidation observed here is representative, our ocean iodine cycling model predicts that future changes in marine nitrification could alter global sea surface I^- fields with potential implications for atmospheric chemistry and air quality.

Introduction

Iodine plays an important role in catalytic ozone destruction and new particle formation in the troposphere, thereby impacting the oxidative capacity of the atmosphere (Sherwen *et al*., 2016) and the Earth's radiation balance (O'Dowd *et al*., 2002). Sea-to-air iodine transfer is known to be the main source of iodine to the atmosphere (Carpenter, 2003; Sherwen *et al*., 2016). Reactive inorganic iodine (I_2, HOI) emissions resulting from the reaction of gas-phase ozone with sea surface iodide (I^-) is now thought to be the dominant mechanism mediating sea-air iodine emissions (Carpenter *et al*., 2013). The strength of the surface reactive iodine flux is related to sea surface I^- concentrations (Carpenter *et al*., 2013) so knowledge of ocean I^- distributions is required in order to estimate the significance of this process. Furthermore, a detailed understanding of the processes controlling inorganic iodine speciation is needed to allow us to develop predictive capacity regarding sea surface I^-, ozone-deposition rates and sea-air emission of reactive iodine.

Total inorganic iodine is found at 400-500 nM in seawater and predominantly exists as iodate (IO_3^-) and I^- (Chance *et al*., 2014) with inter-conversion between these two species alongside physical mixing being the main causes of spatial and temporal variability in sea surface I^-. Iodate is the thermodynamically stable form and the dominant form in the deep ocean. The existence of relatively higher levels of I^- in the euphotic zone (reviewed by Chance *et al*., 2014) has led to the suggestion that IO_3^- reduction to I^- is linked to primary productivity. This theory has been supported by observations of I^- production in cultures of a wide range of marine phytoplankton (e.g. Chance *et al*., 2007; Bluhm *et al*., 2010; Hepach et al., 2020) and some field studies (Chance *et al*., 2010). The mechanism of biogenic iodate reduction to iodide is not yet known, but may be related to senescence processes (Bluhm et al., 2010; Hepach et al., 2020; Carrano et al., 2020). Reduction of IO_3^- to I^- by phytoplankton nitrate reductase enzymes (Hung *et al*., 2005), or macroalgal cell surface reductases (Carrano et al., 2020), has also been suggested but neither has been confirmed as a significant route of conversion.

77

78 Oxidation of I^- back to IO_3^- is the dominant sink for I-, but is a relatively slow reaction with rate
79 estimates ranging from ~4 to 670 nM yr^{-1} (Chance *et al*., 2014; Hardisty *et al*., 2020). The rates and
80 processes involved in I^- to IO_3^- oxidation are associated with large uncertainty (Truesdale *et al*., 2001;
81 Amachi *et al*., 2008), and the mechanisms involved remain undefined. This uncertainty has been
82 suggested to be one of the factors hindering the development of mathematical models of iodine
83 transformations in the global oceans (Truesdale *et al*., 2001). Abiotic oxidation of I^- back to IO_3^- in
84 the ocean (e.g. by oxygen, hydroxyl radicals, hydrogen peroxide and ozone) is thought to occur so
85 slowly as to be insignificant (e.g. Wong, 1991), and so I^- oxidation to IO_3^- is also thought to be
86 associated with marine microbiological activity. I^- oxidation to I_2 has been observed in bacterial
87 isolates obtained from a range of environments including seawater aquaria (Gozlan *et al*., 1968),
88 natural gas brines (Iino *et al*., 2016) and seawater/marine mud (Fuse *et al*., 2003). Additionally,
89 based on field observations, a number of studies (Truesdale *et al*., 2001; Žic *et al*., 2013) have
90 proposed that I^- oxidation to IO_3^- is linked to nitrification in marine systems. Nitrification is the two-
91 stage biological transformation of ammonia (NH_3) to nitrate (NO_3^-) (Equations 1 and 2; Koops &
92 Pommerening-Röser, 2001) mediated by chemoautotrophic ammonia-oxidising bacteria (AOB), and
93 nitrite-oxidising bacteria (NOB). Previously thought to only occur outside of the euphotic zone,
94 nitrification is now known to occur throughout the oceanic water-column (reviewed by Yool *et al*.,
95 2007).

96 $$2NH_4^+ + 3O_2 \rightarrow 2NO_2^- + 4H^+ + 2H_2O \quad (1)$$

97 $$2NO_2^- + O_2 \rightarrow 2NO_3^- \quad (2)$$

98 A link between I^- oxidation/ IO_3^- production and nitrification is yet to be confirmed but, if established,
99 would suggest that I^- oxidation to IO_3^- is widespread throughout the world's oceans (Yool *et al*.,
100 2007).

The primary aim of this study was to establish whether I^- oxidation to IO_3^- is associated with marine nitrification. Our objectives were to determine if IO_3^- production occurs in cultures of marine ammonia- and nitrite-oxidising bacteria supplied with I^-, determine the relative rates of IO_3^- production and nitrification and explore the possible implications of the findings.

Methods

Cultures

Stock bacterial cultures were taken from the existing culture collections of the authors. Two marine AOB cultures, *Nitrosomonas* sp. Nm51 (C-15) and *Nitrosococcus oceani* Nc10 (C-107, ATCC 19707) were investigated for IO_3^- production in the presence of I^- as the only iodine source. These strains were originally isolated from seawater in the south Pacific and the north Atlantic respectively (Watson and Mandel, 1971). Cultures were grown in the dark in a water bath at 25 °C in autoclaved ESAW artificial seawater mixture (Berges *et al*., 2001) made up using distilled water. The ESAW media was supplemented with 7-8 mM ammonium chloride and potassium phosphate. We also conducted preliminary tests on three active marine NOB: *Nitrospira marina* Nb-295 (isolated from Gulf of Maine, Watson et al., 1986); *Nitrospina gracilis* 3/211 (isolated from the south Atlantic, Watson and Waterbury, 1971); *Nitrococcus mobilis* Nb-231 (ATCC 25380, isolated from Galapagos seawater, Watson and Waterbury, 1971). However we saw no evidence of IO_3^- production in any of the NOB cultures studied and these results are not discussed further. Handling of cultures was done at all times in a biosafety cabinet using sterile equipment.

Experimental Set Up

For the AOB experiments triplicate cultures were incubated alongside triplicate media-only controls for periods of 8-12 days. The experiments were kept as short as possible to avoid significant changes in pH in the bulk media which would impact inorganic iodine speciation. Hence experiments were only run until an increase in nitrite across two time-points was observed. Samples were taken at regular intervals of between 1 to 6 days for pH measurement, cell counts and determination of NO_2^-, IO_3^-, I^- and NH_4^+/NH_3 concentrations. In all cases, I^- (Aristar) was added to be at similar concentrations with the NH_4^+ required in the growth media. The levels of I^- are much higher than those encountered in the oceans (global ocean median=77 nM I^- [interquartile range 28-140 nM], Chance *et al*., 2014) but were chosen to be similar to the levels of NH_4^+. This is because in the marine environment nitrifiers would be exposed to similar ratio of NH_4^+ and I^-. For example, Rees *et al*. (2006) show that NH_4^+/NH_3 occurs at concentrations ranging from 60-300 nM in the Atlantic between 60°N to 50°S.

pH

A spectrophotometric method using a Lambda 25 UV/Vis spectrophotometer (Perkin-Elmer) and m-cresol purple dye (Dickson *et al*., 2007) with measurements at 730, 578 and 434 nm was used to determine pH in the cultures and media-only controls. Salinity, needed for the pH calculation, was calculated from conductivity measured using a calibrated Hanna Instruments hand-held probe.

Cell counts

Immediately after sampling, 4 mL of the culture was fixed with 15 μL of 50% glutaraldehyde (Alfa Aesar), flash frozen in liquid nitrogen and placed in a -80 °C freezer for later determination of cell density. Cell counts were made using a Beckman Coulter Cytoflex S flow cytometer (flow rate of 10 μL min^{-1}) within 2 months of collection. DAPI (Sigma; 2 μg mL^{-1}) stained samples were excited by a

147 laser at 405 nm and the emitted fluorescence detected using an avalanche photodiode detector with a
148 reflective band pass filter 450/45. The flow cytometer thresholds were set using the 405 nm laser
149 side scatter and the DAPI fluorescence signals.
150
151 ***Nitrite concentration***
152 NO_2^- was measured in 0.45 μm (Millex) filtered samples using a spectrophotometric method
153 (Lambda 25 UV/Vis spectrophotometer, Perkin-Elmer) developed by Norwitz & Keliher (1984). The
154 method involves diazotizing nitrite with sulfanilamide (Fisher, analytical reagent grade) and coupling
155 with N-1-naphthylethylenediamine dihydrochloride (Fisher, analytical reagent grade) to form a
156 coloured azo dye which is measured spectrophotometrically at 540 nm. The method was calibrated
157 using $NaNO_2$ standards (Fisher, analytical reagent grade) prepared in the ESAW-based media.
158
159 ***Iodate Concentration***
160 IO_3^- concentrations were measured in 0.45 μm (Millex) filtered samples using a manual version of the
161 spectrophotometric (Lambda 25 UV/Vis spectrophotometer) method detailed in Truesdale &
162 Spencer, 1974 and Jickells *et al.*, 1988. Absorbance was measured at 350 nm. Strictly, this method
163 determines all oxidised (0 to +5 oxidation state) forms of inorganic iodine, but in seawater derived
164 media this is predominantly IO_3^-, and so will be referred to as IO_3^- iodate hereafter. The method was
165 calibrated using potassium iodate (Aristar) standard solutions made up in ESAW.
166 Some validation and modification to the method was required due to the nature of our experimental
167 set-up. Chapman & Liss (1977) show that NO_2^- can interfere with spectrophotometric IO_3^-
168 measurements (using sulfamic acid) at ambient seawater concentrations with a 15% error. Clearly
169 significant interference would be an issue for our experiments where NO_2^- was being produced so we
170 ran tests. We found that the presence of NO_2^- up to 10 μM had negligible impact on IO_3^- measurements

(between 0.1-50 μM). We did however identify that the high starting concentration of I^- (~10 μM) in the culture media was problematic. The iodate analysis method comprises two steps: the first involves an initial absorbance reading after the addition of sulfamic acid; the second involves the addition of excess I^-. Under acidic conditions I^- reacts with IO_3^- to form I_2 (equation 3a) which reacts with excess I^- to form the coloured ion I_3^- (equation 3b) that can be measured spectrophotometrically.

$$IO_3^- + 5I^- + 6H^+ \rightarrow 3I_2 + 3H_20 \quad (3a)$$

$$I_2 + I^- \rightarrow I_3^- \quad (3b)$$

The difference between the first and second absorbance readings is then used to calibrate the method. In the case of our experiments the media already contained excess I^- so the formation of I_2 and I_3^- was initiated as soon as the acid was added in the first step. Hence we calibrated the method based on a single absorbance reading obtained after acid and then additional I^- was added. Calibrations and standard checks revealed this approach did not have any impact on the quality of the data.

Ammonium Concentration

NH_4^+ concentrations were measured in 0.45 μm (Millex) filtered samples with a Seal Analytical Autoanalyser 3 according to method G-109-93 rev. 10 (Seal Analytical) using sodium salicylate, dichloro-isocyanuric acid and citrate buffer. The method was calibrated using standards ranging from 0-2 mg/L prepared from dilutions of a 1000 mg/L ammonium standard solution (Merck).

Iodide Concentration

I^- concentrations were determined using a Dionex ICS-2000 ion chromatograph equipped with an EGC III KOH elugen cartridge, AG18 (2 x 50 mm) guard column, AS18 (2 x 250 mm) analytical column, ASRS 300 (2 mm) suppressor, DS6 heated conductivity cell and AS40 autosampler.

194 Samples were diluted 100-fold with 18 MΩ deionised water for analysis and 5 μL was injected onto
195 the ion chromatograph. Aqueous potassium hydroxide was used as the eluent at a flow rate of 0.25
196 mL min^{-1} with a gradient program starting from an initial concentration of 2 mM hydroxide (hold 1
197 min) to 20 mM at 18 min then to 41 mM at 19 min (hold 2 min) before returning to 2 mM. The I^-
198 retention time was 19 min. The instrument was calibrated with matrix-matched standards ranging
199 from 0-100 nM (I^-), prepared from dilutions of a 1000 mg/L iodide standard solution (Fisher
200 Scientific) with 18 MΩ deionised water and containing a final concentration of 1% ESAW.

201

202 ***Data Analysis***

203 As in Guerrero and Jones (1996), the NH_4^+ oxidation rate is defined here as the rate of increase in
204 NO_2^-. Similarly, we define the rate of I^- oxidation as the rate of increase in IO_3^-. This is appropriate as
205 no other iodine species were supplied to the cultures and conversion between I^- and IO_3^- is known to
206 be the main cause of variability in inorganic iodine speciation (Bluhm *et al*., 2010; Chance *et al*.,
207 2014). Average NO_2^- and IO_3^- production rates were calculated for each replicate culture using
208 Equation 4.

209 $$\text{Production Rate (nM day}^{-1}) = \frac{(C_{end} - C_0)}{t} \qquad (4)$$

210 where C_0 and C_{end} are the NO_2^- or IO_3^- concentrations observed at the start and end of the experiment
211 and t is the experimental duration in days. Cell-normalised rates were calculated by dividing these
212 rates by the final cell density observed in each AOB culture and are hence likely to be minimum
213 values.

214

215

216 **Results**

Cell counts and pH

Increases in cell density were observed in all replicates of *Nitrosomonas* sp. and *Nitrosococcus oceani* between the start and end of the experiment indicating growth (Figure 1). Average initial cell density in the *Nitrosomonas* sp. cultures was 21,767 (±4,046) cells mL^{-1} and this increased to 150,983 (±7,585) cells mL^{-1} by the end of the experiment (8 days). For *Nitrosococcus oceani* start and end (12 days) cell densities were 16,947 (± 3,098) and 71,430 (±9,062) cells mL^{-1}, respectively. Average pH levels in the culture experiments calculated from measurements at each time point (data not shown) were 7.69 (±0.07) for *Nitrosomonas* sp. and 7.41 (±0.12) for *Nitrosococcus* sp. These pH levels are consistent with those found in the media-only controls (7.64±0.07 for *Nitrosomonas* sp; 7.64±0.15 for *Nitrosococcus oceani*).

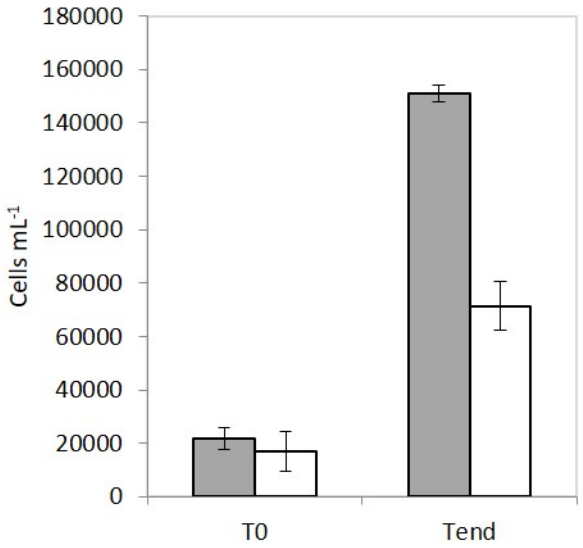

Figure 1. Average cell number in the *Nitrosomonas* sp. (grey bars) and *Nitrosococcus oceani* (white bars) cultures used in this study at the start (T_0) and end (T_{end}; 8 days for *Nitrosomonas* sp. and 12 days for *Nitrosococcus oceani*) of each experiment. Error bars are standard deviations from three replicate cultures.

Iodine and nitrogen speciation

Figure 2 shows that significant increases in the concentrations of IO_3^- (compared to media-only controls) were observed alongside NO_2^- production in both AOB cultures studied. In *Nitrosomonas* sp. (Figure 2ai and 2bi) there was a steady increase in IO_3^- concentrations throughout the experiment reaching a maximum of 19,921 (±4,754) nM by the end of the experiment (day 8). In contrast NO_2^- concentrations reached a maximum of 2,360 (±386) nM by day 6 and remained at around that level until the end of the experiment. In *Nitrosococcus oceani* (Figure 2aii and 2bii) IO_3^- concentrations increased rapidly during the initial stages of the experiment reaching 23, 943 (±8,568) nM by day 6. IO_3^- concentrations at the end of the experiment (day 12) were 16,365 (±7,603) nM. NO_2^- concentrations increased gradually throughout the experiment reaching 5,547 (±1,251) nM by day 12. There was larger variability in IO_3^- concentrations between replicates for *Nitrosococcus oceani* but despite this a clear increase in all replicates was observed.

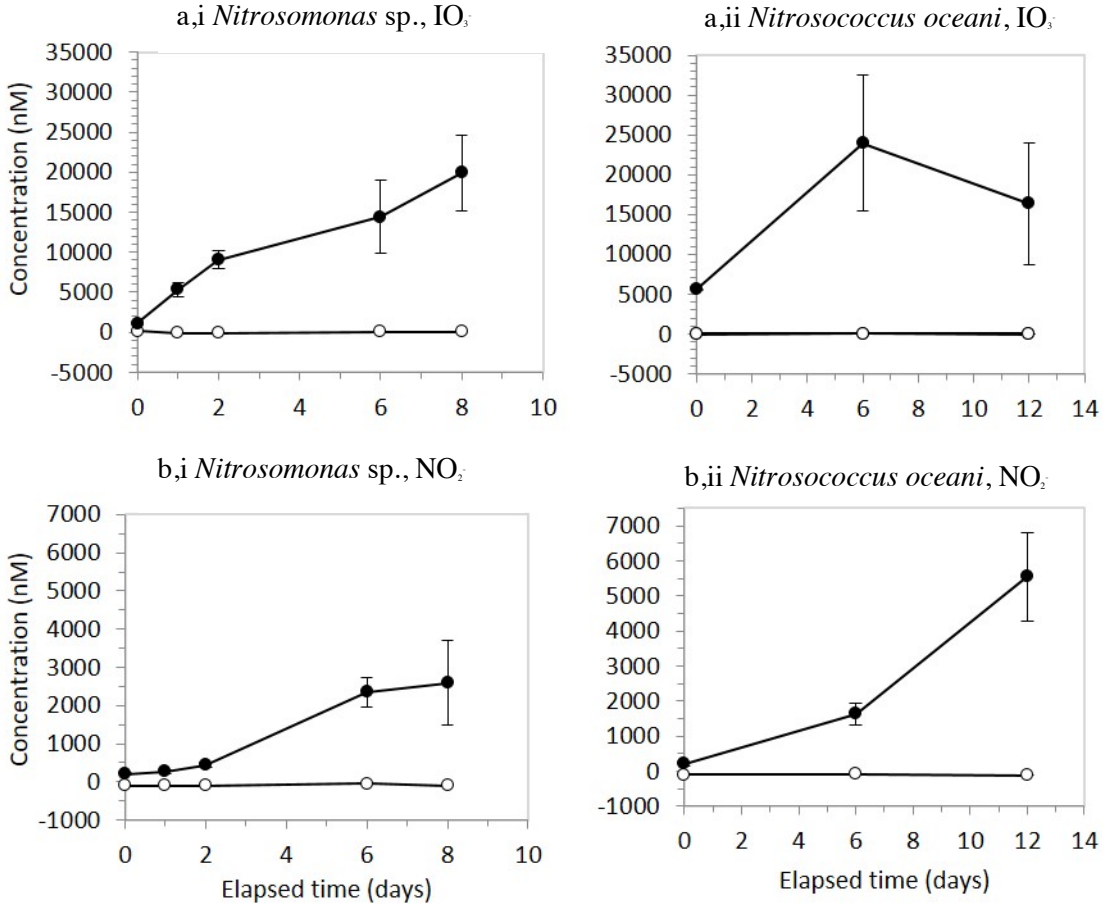

Figure 2. Changes in iodate (a) and nitrite (b) concentrations in cultures (closed symbols) and media-only controls (open symbols) for two cultures of ammonia-oxidising bacteria: i) *Nitrosomonas* sp.; and, ii) *Nitrosococcus oceani* supplied with 9-10 mM iodide and 7-8 mM NH_4^+. Error bars show the standard deviation of three replicate cultures.

Average production rates of IO_3^- and NO_2^- are presented in Table 1. In *Nitrosomonas* sp. average rates (±standard deviation) were 2,348 (±593) nM IO_3^- day^{-1} and 298 (±141) nM NO_2^- day^{-1}. In *Nitrosococcus oceani* averages rates were 897 (±640) nM IO_3^- day^{-1} and 445 (±99) nM NO_2^- day^{-1}. Minimum cell-normalised rates (based on the final cell density observed in each culture) were 15.69 (±4.71) fmol IO_3^- $cell^{-1}$ day^{-1} and 1.96 (±0.88) fmol NO_2^- $cell^{-1}$ day^{-1} for *Nitrosomonas* sp., and 11.96 (±6.96) fmol IO_3^- $cell^{-1}$ day^{-1} and 6.19 (±0.56) fmol NO_2^- $cell^{-1}$ day^{-1} for *Nitrosococcus oceani*. Molar ratios of iodate-to-nitrite production were 9.2±4.0 for *Nitrosomonas* sp. and 1.88±0.91 for *Nitrosococcus oceani*.

Table 1. Nitrite and iodate production rates (± standard deviations) observed in cultures of the ammonia-oxidising bacteria *Nitrosomonas* sp. and *Nitrosococcus oceani*. Cell-normalised values are a minimum as they are calculated using maximum cell densities.

	Nitrite		Iodate	
Culture	**nM day^{-1}**	**fmol $cell^{-1}$ day^{-1}**	**nM day^{-1}**	**fmol $cell^{-1}$ day^{-1}**
Nitrosomonas sp.	298 (±141)	1.96 (±0.88)	2,348 (±593)	15.69 (±4.71)
Nitrosococcus oceani	445 (±99)	6.19 (±0.56)	897 (±640)	11.96 (±6.96)

Figure 3 shows that, within error, a decline in I^- or NH_4^+ concentrations was not observed during either of the AOB experiments. Average starting I^- or NH_4^+ concentrations in *Nitrosomonas* sp. were 9.8 (±0.2) mM and 7.6 (±0.1) mM respectively. At the end of the experiment these values were 10.2 (±0.3) mM I^- and 7.7 (±0.1) mM NH_4^+. For *Nitrosococcus oceani* the start and end concentrations were 9.8 (±0.3) and 9.4 (±0.1) mM for I^- and 7.8 (±0.1) and 7.7 (±0.1) mM for NH_4^+. This result was expected as the average standard deviations associated with the observed concentrations of I^- or NH_4^+ (i.e. 0.1 to 0.3 mM) are at least an order of magnitude higher than the maximum levels of IO_3^- and

NO_2^- observed in the culture experiments, i.e. very little of the initial stock of NO_2^- or NH_4^+ was oxidised during the experiments.

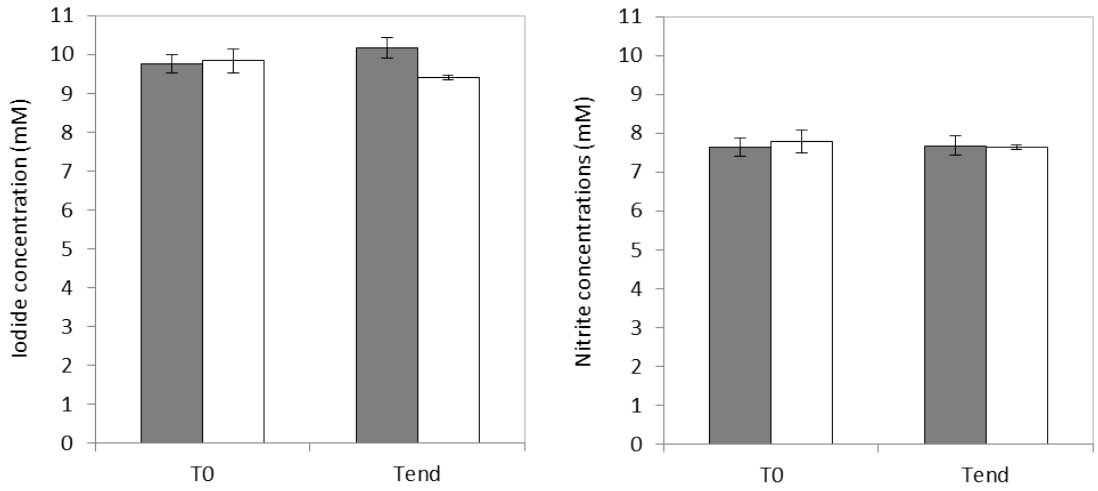

Figure 3. Start and end concentrations of a) iodide and b) ammonia in cultures of *Nitrosomonas* sp. (grey bars) and *Nitrosococcus oceani* (white bars). Error bars show the standard deviation of three replicate cultures.

Discussion

Iodate production by ammonia-oxidising bacteria

Our results confirm that IO_3^- production occurs in cultures of the ammonia-oxidising bacteria *Nitrosomonas* sp. and *Nitrosococcus oceani* supplied with I^-, but not in cultures of nitrite oxidising bacteria. Coincident increases in NO_2^- (Figure 2) show that both cultures were actively oxidising ammonia throughout the experiments at rates of 1.96±0.0.88 fmol NO_2^- $cell^{-1}$ day^{-1} for *Nitrosomonas* sp. and 6.19±0.56 fmol NO_2^- $cell^{-1}$ day^{-1} for *Nitrosococcus oceani*. Whilst these cell-normalised oxidation rates are of the same order as those reported in the literature (e.g. 6-20 fmol NO_2^- $cell^{-1}$ day^{-1}; Ward *et al*., 1987; 1989) they are at the lower end. This is consistent with the approach taken here to calculate the rates by normalising to the final (highest) cell densities. It is also worth noting that the

297 cultures were at an early stage of growth and had relatively low cell densities during the experiment.
298 This was done to avoid significant changes in pH in the bulk media which would impact inorganic
299 iodine speciation (*Section 3.2*). The observation of an increase in IO_3^- concentrations alongside active
300 biological ammonia oxidation supports previous studies (e.g. Truesdale *et al*., 2001; Zic *et al*., 2013)
301 which have shown that high aqueous concentrations of IO_3^- are found in regions of enhanced
302 nitrification, and provides the first direct confirmation of a biological basis for at least one
303 mechanism of iodide oxidation
304
305 Whilst we did not set out to establish the mechanism for I^- to IO_3^- oxidation by marine nitrifiers, some
306 speculations can be made. As I^- oxidation to IO_3^- requires the transfer of six electrons, it may occur in
307 a series of one- or two- electron transfer steps. Initially, I^- may be oxidised to molecular iodine ($I^- \rightarrow$
308 I_2), a reaction which is thermodynamically unfavourable at the pH of seawater (Luther *et al*., 1995).
309 Further oxidation to IO_3^- by disproportionation ($I_2 \rightarrow HOI \rightarrow IO_3^-$) can occur spontaneously, but in
310 seawater is subject to competition with reduction of I_2 by organic matter (Truesdale & Moore, 1992;
311 Truesdale *et al*., 1995). It is not known whether the ammonia-oxidisers mediate just the first stage of
312 I^- oxidation, with the observed IO_3^- production due to subsequent spontaneous reactions in the culture
313 media, or if they are involved in driving the complete conversion of I^- to IO_3^-. However, bacteria
314 which just oxidise I^- to I_2 have been isolated from seawater aquaria (Gozlan, 1968), I-rich natural gas
315 brine waters (Amachi *et al*., 2005) and marine environmental samples (Fuse *et al*., 2003; Amachi *et*
316 *al*., 2005).
317
318 The observed IO_3^- production is either linked to the nitrification process itself or associated with other
319 metabolic activities of the AOB studied. Truesdale *et al*. (2001) has proposed that I^- oxidation to IO_3^-
320 would be energetically advantageous for chemoautotrophic AOB. In that case the key enzymes used
321 to obtain energy during the oxidation of NH_4^+ to NO_2^- (ammonia monooxygenase [AMO] and

hydroxylamine oxidoreductase [HAO]) could also have the potential to use I^- as a substrate. The observed IO_3^--to-NO_2^- molar production rates (9.2±4.0 for *Nitrosomonas* sp. and 2.3±1.1 for *Nitrosococcus oceani*) are intriguing. If AMO/HAO are involved, this suggests that the enzymes have higher affinities for I^- than NH_4^+/NH_2OH given the similar concentrations of I^- and NH_4^+ used in the experiments. Other enzymes that have been implicated in I^- oxidation include the chloroperoxidases (Thomas & Hager, 1968) but we do not know if they occur in AOB. The exact metabolic pathway driving the observed IO_3^- production and its controls (i.e. substrate concentrations, light intensity) will need to be determined in future work. To establish if such further experimentation is warranted we need to explore whether the link between nitrification and I^- oxidation is likely to be an important part of inorganic iodine cycling in seawater.

Implications for inorganic iodine speciation in the oceans

Our culture studies suggest that the molar rate of I^- oxidation (IO_3^- production) is ~2-9 times higher than that for ammonia oxidation (nitrification). Note that although ammonium and iodide concentrations were much higher in the experimental media than in the oceans, the concentration ratio of these species was comparable to that found naturally. Ammonia oxidation rates in seawater range from below detection to 10^2 nM day^{-1} (Table 2). Literature estimates of the rate of I^- oxidation in the marine environment range from ~4 to 670 nM $year^{-1}$ or 0.01 to 1.84 nM day^{-1} (reviewed in Chance *et al*., 2014). If the oxidation molar ratios observed in this study (~2-9) are representative, predicted rates of I^- oxidation are in-line (i.e. 2-9 times higher) with the lower end of observed ammonia oxidation rates (Table 2).

Table 2. Ammonia-oxidation rates measured in a range of ocean regions.

Study	Location	Rate (nM day^{-1})
Newell *et al*. (2011)	Arabian Sea, Indian Ocean	undetected to 21.6
Smith *et al*. (2015)	Northeast Pacific	< 0.01 to 90
Peng *et al*. (2015)	Eastern tropical north Pacific	< 1to 8.6

Newell *et al*. (2013)	Subtropical Atlantic, Sargasso Sea (BATS)	< 2
Lam *et al*. (2007)	Black Sea	7-24
Beman *et al*. (2012)	Gulf of California, eastern tropical north Pacific	0-348

346

347

348 Truesdale *et al*. (2001) derive likely I^- oxidation (or IO_3^- production) rates for the near surface Black
349 Sea using an iodine budget and this allows us to examine the potential importance of the link
350 between nitrification and I^- oxidation on a local scale. They predict a minimum I^- oxidation flux of
351 3.89 x 10^{-4} mol I m^{-2} $year^{-1}$ which is an average of 0.02 nM day^{-1} at a mixed-layer depth (MLD) of 50 m
352 or 0.11 nM day^{-1} at an MLD of 10 m. Lam *et al*. (2007) report an AOB abundance of ≤1,400 cells
353 mL^{-1} in the Black Sea. If we apply a cell density of 1,400 AOB cells mL^{-1} to the average cell-
354 normalised rates of IO_3^- production observed in this study (Table 1) we derive I^- oxidation rates of
355 ~20 nM d^{-1}. This is clearly much higher than the rates suggested in Truesdale *et al*. (2001). This
356 discrepancy could be explained in a number of ways. Firstly, Lam *et al*. (2007) state that net
357 nitrification only takes place within a narrow depth range of the Black Sea water column (i.e.
358 between 71 and 81 m) and, the I^- oxidation values derived in Truesdale *et al*. (2001) are minimum
359 values. It is also possible that the AOB studied here have a higher capacity for I^- oxidation (per unit
360 ammonia-oxidised) than other ammonia-oxidisers or that our culture conditions (e.g. substrate
361 availability) promoted higher I^- oxidation rates than would be observed in marine systems. For
362 example, ammonia-oxidising Archaea (AOA), which can outnumber known bacterial ammonia
363 oxidisers by orders of magnitudes in environments such as the marine water-column (reviewed by
364 Schleper & Nicol, 2010), may have a very different capacity for I^- oxidation compared to the AOB
365 studied here. Further studies are needed to establish the relationship between ammonia- and I^-
366 oxidation in the marine environment.

367

368 ***Potential implications for future oceanic inorganic iodine distributions***

Environmental factors which are known to be currently undergoing change in the oceans (e.g. oxygen, light, pH, temperature) have all been found to impact rates and patterns of marine nitrification (reviewed by Pajores and Ramos, 2019). Whilst there remains some uncertainty about the future magnitude and, in some cases, sign of the response, some of the expected future changes in marine nitrification are large. For example, whilst some studies have seen no impact on specific marine nitrifiers (e.g. Qin et al., 2014), Beman et al. (2011) suggest that expected rates of acidification could cause a decline in ammonia oxidation by up to 44% within the next few decades. It is hence worth exploring how possible future changes in marine nitrification could impact ocean iodine cycling.

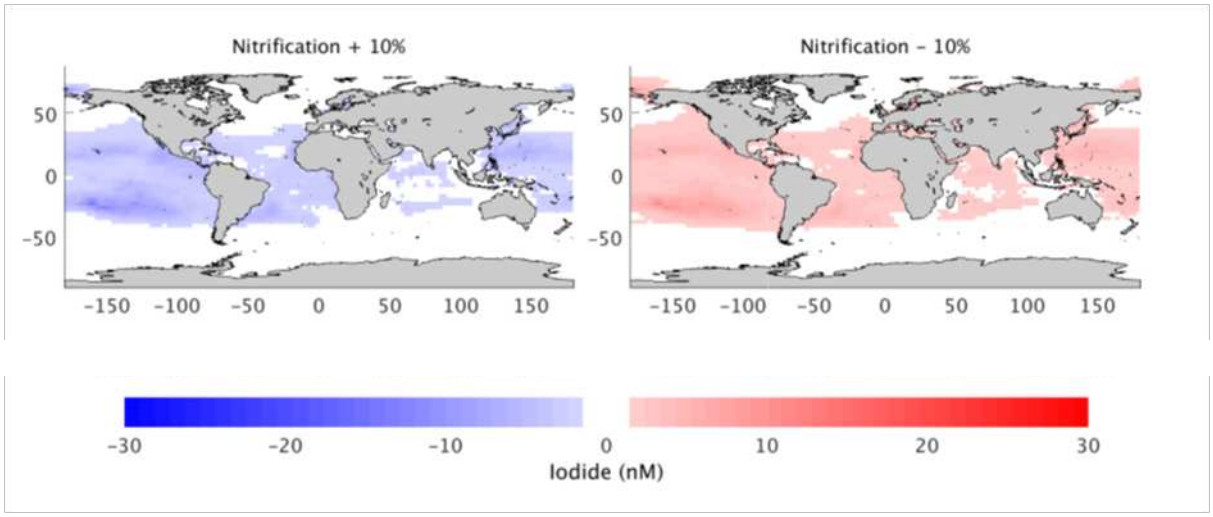

Figure 4. Modelled changes in surface I^- concentration (nM) resulting from a) +10%, b) -10%, changes in the rates of nitrification. Negative percent values indicate a decline in the rate of nitrification and *vice-versa*. Negative values on the scale bar indicate a decrease in I^- concentrations and *vice versa*.

In order to explore the possible impact of future changes in marine nitrification rates on sea surface iodine fields we used the ocean cycling model described in Wadley et al. (2020). Within the model iodide production is driven by primary productivity, and I^- oxidation to IO_3^- linked to nitrification in the mixed layer. Nitrogen fluxes and the spatial distribution of mixed layer ammonia oxidation are derived from a global biogeochemical cycling model (Yool *et al.*, 2007). I^- is oxidised to IO_3^- in

390 association with the ammonia oxidation, with the same I:N:C ratio as associated with iodide
391 production (Truesdale *et al*., 2001; Long *et al*., 2015). The model does not use any of the rates
392 derived in the current study as these are based on results from only 2 AOB species cultured at high
393 substrate concentrations. Model outputs (Figure 4) show that even with small (+/- 10%) changes in
394 ammonia oxidation there is a clear alteration to sea surface I^- fields. Sea surface I^- concentrations
395 increase as ammonium oxidation rates decrease and *vice-versa*. For example, the ocean cycling
396 model suggests there could be an average global increase of 0.13 nM I^- per 1% decrease in
397 nitrification. The outputs suggest that the change in the iodine fields is spatially variable and will
398 increase as the perturbation to ammonia oxidation increases. For example, at the 44% decline in
399 nitrification predicted by Beman et al. (2011) the model predicts there will be a 25% increase (+30
400 nM) in sea surface I^- in the sub-tropical gyres. Carpenter *et al*. (2013) show that I_2 emissions due to
401 ozone deposition increase near linearly with I^- concentration. Hence, the predicted changes to sea
402 surface I^- fields under future ocean acidification could have a major impact on ozone deposition to
403 the sea surface, atmospheric chemistry and resulting sea-air iodine emissions.

404

405

406 **5.3.Conclusions**

407 This study has shown that I^- oxidation to IO_3^- occurs in cultures of ammonia oxidising (nitrifying)
408 bacteria, but not nitrite oxidising bacteria. Our calculations suggest that I^- oxidation by AOB could be
409 an important control on inorganic iodine speciation in seawater, but to confirm this further study is
410 needed on a wider range of ammonia-oxidisers including ammonia oxidising archaea (AOA).
411 Simulations from our iodine cycling model suggest that changes in nitrification rate, such as those
412 predicted to occur under acidification (Beman *et al*., 2011), could have an important impact on sea
413 surface I^- fields. A future change in marine nitrification could alter sea surface I^- fields. In turn, this

414 could lead to a change in ozone deposition to the sea surface and sea-air iodine emissions with
415 potentially major implications for atmospheric chemistry and air quality.
416
417
418
419
420 **Acknowledgements**

421 We thank Eva Spieck (University of Hamburg, Germany) for supplying the nitrite oxidising bacteria
422 (NOB) cultures used in our preliminary experiments.

423

424 **Author contributions**

425 Claire Hughes: Conceptualisation, Methodology, Formal Analysis, Investigation, Writing – Original
426 Draft, Writing – Review & Editing, Visualisation, Supervision, Project Administration, Funding
427 Acquisition. Eleanor Barton: Formal Analysis, Investigation, Writing – Original Draft. Helmke
428 Hepach: Methodology, Validation, Investigation, Resources. Rosie Chance: Conceptualisation,
429 Methodology, Resources, Data Curation, Writing – Review & Editing, Project Administration,
430 Funding Acquisition. Matt Pickering: Methodology, Validation, Resources. Karen Hogg:
431 Methodology, Resources. Andreas Pommerening-Röser: Methodology, Resources. Martin R.
432 Wadley: Conceptualisation, Methodology, Software, Validation, Formal Analysis, Investigation,
433 Writing – Original Draft, Visualisation. David P. Stevens: Conceptualisation, Methodology,
434 Resources, Supervision, Funding Acquisition. Tim D. Jickells: Conceptualisation, Methodology,
435 Supervision, Funding Acquisition.

436

Data Statement: The data presented in this manuscript is available from the British Oceanographic Data Centre (BODC), doi:10.5285/aaa4dab8-7c9d-0cca-e053-6c86abc00b1d

Funding: This work was supported by the Natural Environment Research Council [grant no. NE/N01054X/1]; and the University of York Laidlaw Scholarship [work undertaken by Eleanor Barton].

Declaration of interest: The authors have no financial conflicts of interest with the research in this paper.

454 **References**

455 Amachi, S., 2008. Microbial Contribution to Global Iodine Cycling: Volatilization, Accumulation,
456 Reduction, Oxidation, and Sorption of Iodine. Microbes Environ., 23(4):
457 269.10.1264/jsme2.ME08548

458 Amachi, S., Muramatsu, Y., Akiyama, Y., Miyazaki, K., Yoshiki, S., Hanada, S., Kamagata, Y.,
459 Ban-nai, T., Shinoyama, H. and Fujii, T., 2005. Isolation of iodide-oxidizing bacteria from
460 iodide-rich natural gas brines and seawaters. Microb. Ecol., 49(4): 547

461 Beman, J.M., Chow, C.-E., King, A.L., Feng, Y., Fuhrman, J.A., Andersson, A., Bates, N.R., Popp,
462 B.N. and Hutchins, D.A., 2011. Global declines in oceanic nitrification rates as a
463 consequence of ocean acidification. Proceedings of the National Academy of Sciences,
464 108(1): 208.10.1073/pnas.1011053108

465 Beman, J.M., Popp, B. N. & Alford, S. E., 2012. Quantification of ammonia oxidation rates and
466 ammonia-oxidizing archaea and bacteria at high resolution in the Gulf of California and
467 eastern tropical North Pacific Ocean. Limnology and Oceanography, 57, 712-726

468 Bluhm, K., Croot, P., Wuttig, K. and Lochte, K., 2010. Transformation of iodate to iodide in marine
469 phytoplankton driven by cell senescence. Aquat. Biol., 11(1): 1.10.3354/ab00284

470 Carpenter, L.J., 2003. Iodine in the marine boundary layer. Chem. Rev., 103(12):
471 4953.10.1021/cr0206465

472 Carpenter, L.J., MacDonald, S.M., Shaw, M.D., Kumar, R., Saunders, R.W., Parthipan, R., Wilson,
473 J. and Plane, J.M.C., 2013. Atmospheric iodine levels influenced by sea surface emissions of
474 inorganic iodine. Nature Geoscience, 6(2): 108.10.1038/ngeo1687

475 Carrano, M.W., Yarimizu, K., Gonzales, J.L., Cruz-López, R., Edwards, M.S., Tymon, T.M.,
476 Küpper, F.C. and Carrano, C.J. 2020. The influence of marine algae on iodine speciation in
477 the coastal ocean. Algae, 35 (2): 167-176. 10.4490/algae.2020.35.5.25

478 Chance, R., Baker, A.R., Carpenter, L. and Jickells, T.D., 2014. The distribution of iodide at the sea
479 surface. Environmental Science-Processes & Impacts, 16(8): 1841.10.1039/c4em00139g
480 Chance, R., Malin, G., Jickells, T. and Baker, A.R., 2007. Reduction of iodate to iodide by cold
481 water diatom cultures. Mar. Chem., 105(1-2): 169.10.1016/j.marchem.2006.06.008
482 Chance, R., Weston, K., Baker, A.R., Hughes, C., Malin, G., Carpenter, L., Meredith, M.P., Clarke,
483 A., Jickells, T.D., Mann, P. and Rossetti, H., 2010. Seasonal and interannual variation of
484 dissolved iodine speciation at a coastal Antarctic site. Mar. Chem., 118(3-4):
485 171.10.1016/j.marchem.2009.11.009
486 Chapman, P. and Liss, P.S., 1977. Effect of Nitrite on Spectrophotometric Determination of Iodate in
487 Seawater. Mar. Chem., 5(3): 243
488 Dickson, A. G., Sabine, C. L. & Christian, J. R. (Eds.), 2007. Guide to Best Practices for Ocean CO_2
489 Measurements. PICES Special Publication 3. 191 pp.
490 Fuse, H., Inoue, H., Murakami, K., Takimura, O. and Yamaoka, Y., 2003. Production of free and
491 organic iodine by *Roseovarius* spp. FEMS Microbiol. Lett., 229(2): 189.10.1016/s0378-
492 1097(03)00839-5
493 Gozlan, R.S., 1968. Isolation of Iodine-Producing Bacteria from Aquaria. Antonie Van
494 Leeuwenhoek Journal of Microbiology, 34(2): 226
495 Hardisty DS, Horner TJ, Wankel SD, Blusztajn J, and Nielsen SG. 2020. Experimental observations
496 of marine iodide oxidation using a novel sparge-interface MC-ICP-MS technique. Chem
497 Geol, 532:119360. 10.1016/j.chemgeo.2019.119360.
498 Hepach H, Hughes C, Hogg K, Collings S, and Chance R., 2020. Senescence as the main driver of
499 iodide release from a diverse range of marine phytoplankton. Biogeosciences, 17:2453–71.
500 https://doi.org/10.5194/bg-17-2453-2020.
501 Hung, C.C., Wong, G.T.F. and Dunstan, W.M., 2005. Iodate reduction activity in nitrate reductase
502 extracts from marine phytoplankton. Bull. Mar. Sci., 76(1): 61

503 Iino, T., Ohkuma, M., Kamagata, Y. and Amachi, S., 2016. *Iodidimonas muriae* gen. nov., sp. nov.,
504 an aerobic iodide-oxidizing bacterium isolated from brine of a natural gas and iodine
505 recovery facility, and proposals of *Iodidimonadaceae* fam. nov., *Iodidimonadales* ord. nov.,
506 *Emcibacteraceae* fam. nov. and *Emcibacterales* ord. nov. Int. J. Syst. Evol. Microbiol.,
507 66(12): 5016.https://doi.org/10.1099/ijsem.0.001462
508 Jickells, T.D., Boyd, S.S. and Knap, A.H., 1988. Iodine cycling in the Sargasso Sea and the Bermuda
509 Inshore waters. Mar. Chem., 24(1): 61.10.1016/0304-4203(88)90006-0
510 Koops, H.-P. and Pommerening-Röser, A., 2001. Distribution and ecophysiology of the nitrifying
511 bacteria emphasizing cultured species. FEMS Microbiol. Ecol., 37(1): 1.10.1111/j.1574-
512 6941.2001.tb00847.x
513 Lam, P., Jensen, M.M., Lavik, G., McGinnis, D.F., Müller, B., Schubert, C.J., Amann, R.,
514 Thamdrup, B. and Kuypers, M.M.M., 2007. Linking crenarchaeal and bacterial nitrification
515 to anammox in the Black Sea. Proceedings of the National Academy of Sciences, 104(17):
516 7104.10.1073/pnas.0611081104
517 Li, H.P., Daniel, B., Creeley, D., Grandbois, R., Zhang, S.J., Xu, C., Ho, Y.F., Schwehr, K.A.,
518 Kaplan, D.I., Santschi, P.H., Hansel, C.M. and Yeager, C.M., 2014. Superoxide Production
519 by a Manganese-Oxidizing Bacterium Facilitates Iodide Oxidation. Appl. Environ.
520 Microbiol., 80(9): 2693.10.1128/aem.00400-14
521 Long, A., Dang, A., Xiao, H. and Yu, X., 2015. The Summer Distribution of Dissolved Inorganic
522 Iodine along 18°N in the South China Sea. Journal of Marine Science: Research &
523 Development, 5: 169.10.4172/2155-9910.1000169
524 Newell, S.E., Babbin, A.R., Jayakumar, A. and Ward, B.B., 2011. Ammonia oxidation rates and
525 nitrification in the Arabian Sea. Global Biogeochemical Cycles,
526 25(4).10.1029/2010GB003940

527 Newell, S.E., Fawcett, S.E. and Ward, B.B., 2013. Depth distribution of ammonia oxidation rates and
528 ammonia-oxidizer community composition in the Sargasso Sea. Limnology and
529 Oceanography, 58(4): 1491.10.4319/lo.2013.58.4.1491
530 Norwitz, G. and Keliher, P.N., 1984. Spectrophotometric determination of nitrite with composite
531 reagents containing sulphanilamide, sulphanilic acid or 4-nitroaniline as the diazotisable
532 aromatic amine and N-(1-naphthyl)ethylenediamine as the coupling agent. Analyst, 109(10):
533 1281.10.1039/AN9840901281
534 O'Dowd, C.D., Jimenez, J.L., Bahreini, R., Flagan, R.C., Seinfeld, J.H., Hameri, K., Pirjola, L.,
535 Kulmala, M., Jennings, S.G. and Hoffmann, T., 2002. Marine aerosol formation from
536 biogenic iodine emissions. Nature, 417(6889): 632
537 Pajores, S. and R. Ramos (2019) Processes and microorganisms involved in the marine nitrogen
538 cycle: knowledge and gaps, Frontiers in Marine Science, 6, 10.3389/fmars.2019.00739
539 Peng, X., Fuchsman, C.A., Jayakumar, A., Oleynik, S., Martens-Habbena, W., Devol, A.H. and
540 Ward, B.B., 2015. Ammonia and nitrite oxidation in the Eastern Tropical North Pacific.
541 Global Biogeochemical Cycles, 29(12): 2034.10.1002/2015GB005278
542 Qin, W., Amin, S.A., Martens-Habbena, W., Walker, C.B., Urakawa, H., Devol, A.H., Ingalls, A.E.,
543 Moffett, J.W., Armbrust, E.V. and Stahl, D.A., 2014. Marine ammonia-oxidizing archaeal
544 isolates display obligate mixotrophy and wide ecotypic variation. Proc. Natl. Acad. Sci. U. S.
545 A., 111(34): 12504.10.1073/pnas.1324115111
546 Schleper, C. and Nicol, G.W., 2010. Ammonia-oxidising archaea--physiology, ecology and
547 evolution. Adv. Microb. Physiol., 57: 1.10.1016/b978-0-12-381045-8.00001-1
548 Schulz, K.G., Barcelos e Ramos, J., Zeebe, R.E. and Riebesell, U., 2009. CO_2 perturbation
549 experiments: similarities and differences between dissolved inorganic carbon and total
550 alkalinity manipulations. Biogeosciences, 6(10): 2145.10.5194/bg-6-2145-2009

551 Sherwen, T., Evans, M.J., Carpenter, L.J., Andrews, S.J., Lidster, R.T., Dix, B., Koenig, T.K.,
552 Sinreich, R., Ortega, I., Volkamer, R., Saiz-Lopez, A., Prados-Roman, C., Mahajan, A.S. and
553 Ordonez, C., 2016. Iodine's impact on tropospheric oxidants: a global model study in GEOS-
554 Chem. Atmospheric Chemistry and Physics, 16(2): 1161.10.5194/acp-16-1161-2016
555 Smith, J.M., Damashek, J., Chavez, F.P. and Francis, C.A., 2016. Factors influencing nitrification
556 rates and the abundance and transcriptional activity of ammonia-oxidizing microorganisms in
557 the dark northeast Pacific Ocean. Limnology and Oceanography, 61(2):
558 596.10.1002/lno.10235
559 Thomas, J.A. and Hager, L.P., 1968. The peroxidation of molecular iodine to iodate by
560 chloroperoxidase. Biochemica Biophysica Research Communications, 32: 770
561 Truesdale, V.W. and Luther, G.W., 1995. Molecular iodine reduction by natural and model organic
562 substances in seawater. Aquatic Geochemistry, 1(1): 89
563 Truesdale, V.W. and Moore, R.M., 1992. Further-Studies on the Chemical-Reduction of Molecular-
564 Iodine Added to Seawater. Mar. Chem., 40(3-4): 199
565 Truesdale, V.W. and Spencer, C.P., 1974. Studies on the determination of inorganic iodine in
566 seawater. Mar. Chem., 2(1): 33.https://doi.org/10.1016/0304-4203(74)90004-8
567 Truesdale, V.W., Watts, S.F. and Rendell, A.R., 2001. On the possibility of iodide oxidation in the
568 near-surface of the Black Sea and its implications to iodine in the general ocean. Deep-Sea
569 Research Part I-Oceanographic Research Papers, 48(11): 2397.10.1016/s0967-
570 0637(01)00021-8
571 Wadley, M.R., Stevens, D.P., Jickells, T.D., Hughes, C., Chance, R., Hepach, H., Tinel, L. and
572 Carpenter, L.J., 2020. A Global Model for Iodine Speciation in the Upper Ocean. Global
573 Biogeochemical Cycles, e2019GB006467.10.1029/2019GB006467

574 Ward, B.B., 1987. Nitrogen transformations in the Southern California Bight. Deep Sea Research
575 Part A. Oceanographic Research Papers, 34(5): 785.https://doi.org/10.1016/0198-
576 0149(87)90037-9
577 Ward, B.B., Glover, H.E. and Lipschultz, F., 1989. Chemoautotrophic activity and nitrification in the
578 oxygen minimum zone off Peru. Deep Sea Research Part A. Oceanographic Research Papers,
579 36(7): 1031.https://doi.org/10.1016/0198-0149(89)90076-9
580 Watson, S.W., Bock, E., Valois, F.W., Waterbury, J.B. and Schlosser, U., 1986. *Nitrospira*
581 *marina* gen. nov. sp. nov.: a chemolithotrophic nitrite-oxidizing bacterium. *Arch*.
582 *Microbiol*. 144**:** 1–7. https://doi.org/10.1007/BF00454947
583 Watson S. W. and Mandel , M., 1971. Comparison of the morphology and deoxyribonucleic acid
584 composition of 27 strains of nitrifying bacteria. J. Bacteriol., 107 (2): 563-569
585 Watson, S.W. and Waterbury, J.B., 1971. Characteristics of two marine nitrite oxidizing
586 bacteria, *Nitrospina gracilis nov. gen. nov. sp.* and *Nitrococcus mobilis nov. gen. nov.*
587 *sp. . Archiv. Mikrobiol*. 77**:** 203–230. https://doi.org/10.1007/BF00408114
588 Wong, G.T.F., 1991. The marine geochemistry of iodine. Reviews in Aquatic Sciences, 4(1): 45
589 Yool, A., Martin, A.P., Fernandez, C. and Clark, D.R., 2007. The significance of nitrification for
590 oceanic new production. Nature, 447(7147): 999.10.1038/nature05885
591 Zic, V., Caric, M. and Ciglenecki, I., 2013. The impact of natural water column mixing on iodine and
592 nutrient speciation in a eutrophic anchialine pond (Rogoznica Lake, Croatia). Estuarine
593 Coastal and Shelf Science, 133: 260.10.1016/j.ecss.2013.09.008

Supplementary Data

Iodate production in cultures of marine ammonia-oxidising bacteria: implications for future inorganic iodine distributions in the oceans

Claire Hughes, Eleanor Barton, Helmke Hepach, Rosie Chance, Matt Pickering, Karen Hogg, Andreas Pommerening-Röser, Martin R. Wadley, David P. Stevens & Tim D. Jickells

Table S1: Concentrations of iodine and nitrogen species, cell counts and pH in bacterial cultures and media-only controls over time

		[Iodate] and [Nitrite], nmol L-1																					
			Media-only control			Bacteria			Cell count, cells/mL			[Ammonium], mmol L-1			[Iodide], mmol L-1			pH					
									Bacteria			Bacteria			Bacteria			Media-only control			Bacteria		
Bacterial strain	Day	Analyte	A	B	C	A	B	C	A	B	C	A	B	C	A	B	C	A	B	C	A	B	C
Nitrosomonas sp.	0	*Nitrite:*	0.0	0.0	0.0	0	0	0	18140	26130	21030	7.64	7.64	7.61	9.94	9.49	9.84	7.69	7.7	7.71	7.75	7.77	7.78
	1		-0.7	-2.3	-3.2	31	63	71										7.58	7.65	7.64	7.67	7.63	7.7
	2		2.5	2.3	-2.3	257	210	238															
	6		45.1	44.5	37.5	2175	1790	2488															
	8		2.1	-4.6	-7.2	3485	2430	1228	158610	143440	150900	7.75	7.61	7.68	10.4	10.2	9.9	7.54	7.61	7.61	7.62	7.61	7.65
	0	*Iodate:*	0	0	0	0	0	0															
	1		-98	-98	-212	4047	3529	5022															
	2		-125	-125	-202	8741	6582	8490															
	6		19	19	-58	18112	8764	12893															
	8		35	35	-85	16403	24243	15702															
Nitrosococcus oceani	0	*Nitrite:*	0	0	0	0	0	0	19870	17270	13700	7.83	7.72	7.78	9.5	10.0	10.0	7.61	7.60	7.59	7.54	7.50	7.52
	6		30.3	21.3	23.4	1675	1428	1194										7.68	7.69	7.64	7.45	7.44	7.49
	12		-12.9	-12.6	-10	6705	4718	4604	81830	67230	65230	7.69	7.58	7.68	9.5	9.4	9.4	7.66	7.67	7.63	7.21	7.30	7.27
	0	*Iodate:*	0	0	0	0	0	0															
	6		72	59	61	15000	28036	11970															
	12		25	28	46	19569	7185	5519															